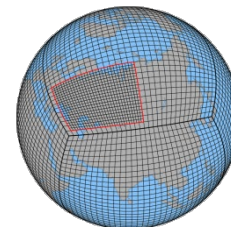


Разработка глобальной модели атмосферы для исследований климата
и прогноза погоды с высоким пространственным разрешением

Development of the global atmospheric model for high-resolution climate
modeling and weather prediction

V.Shashkin, G.Goyman, D.Markhanov, I.Tretyak, Z.Khaidapov

Kislovodsk, 08.04.2026



- Motivation
- Model design and features
- Test-driven development strategy
- **Aquaplanet results with full subgridscale physics**
- A word on computational efficiency
- Conclusions & outlook



- 
- Universal atmospheric dynamics solver for various problems:
 - Numerical weather prediction (with SLAV parameterizations)
 - Climate (with ARTS-ESM/INMCM parameterizations)
 - (Upper atmosphere dynamics)
 - New features (for Russian global atmospheric codes):
 - High-resolution, non-hydrostatic, Local mesh refinement
 - Better software
 - Readable & structured code
 - Flexible (various equation systems)
 - Tests for basic units and combinations
 - **Scalable algorithms & GPU**

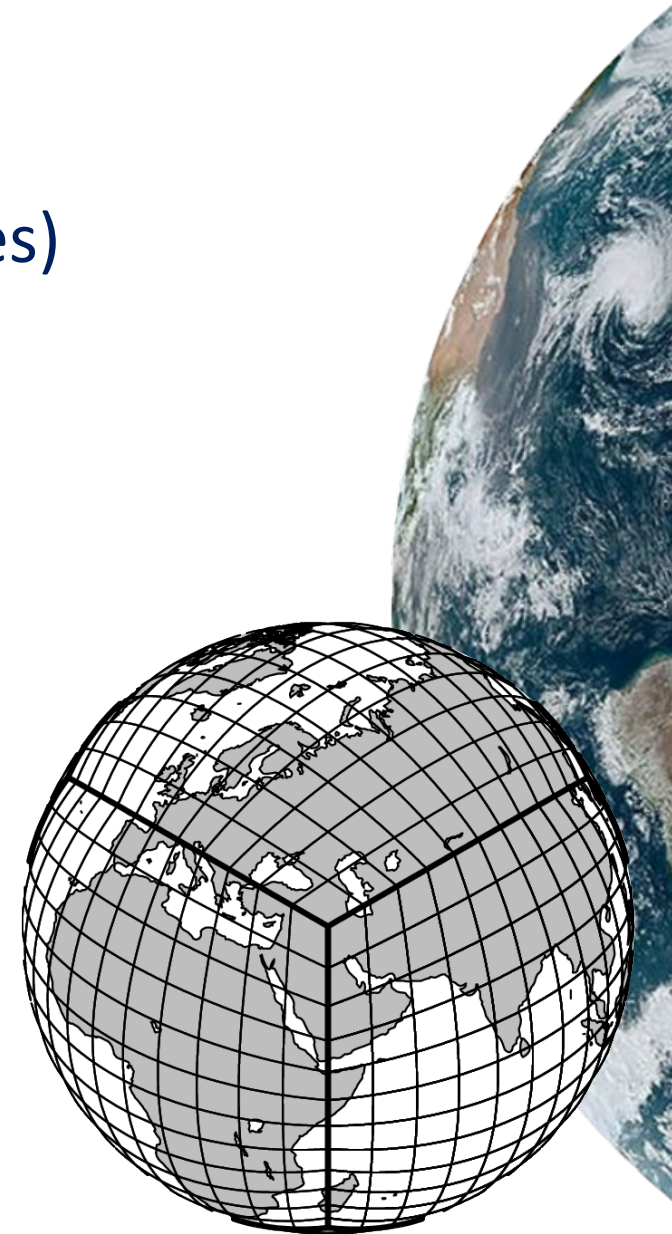
Basic features

- Cubed-sphere grid (quasi-uniform spherical grid)
- Stable and conservative high-order accurate finite-difference method (Summation-by-Parts finite-differences)

Discrete Ostrogradsky-Gauss theorem:

$$\int_{\Omega} (\vec{v} \cdot \nabla p + p \nabla \cdot \vec{v}) d\Omega = 0$$

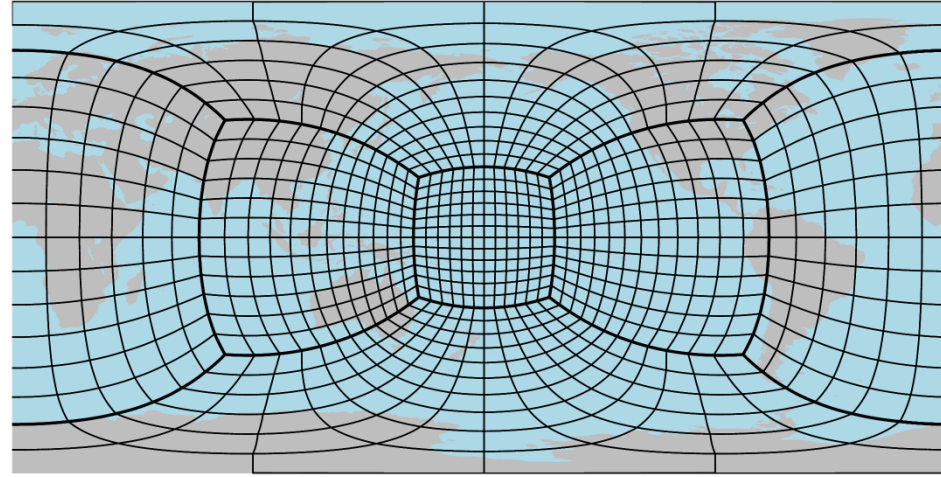
- Hydrostatic & non-hydrostatic versions
- Collection of numerical methods (possibility to create combinations optimal for specific tasks)



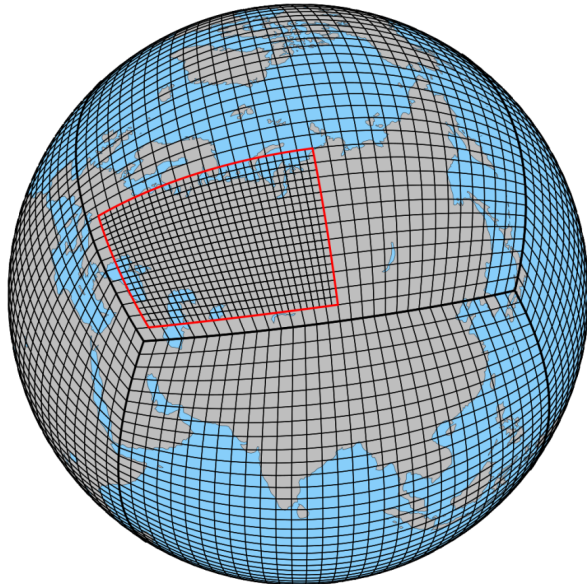
Mesh refinement strategies

1. Schmidt transform

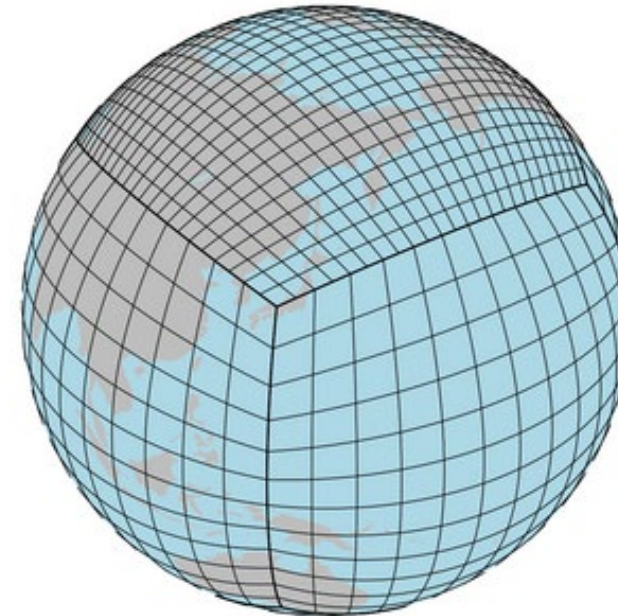
$$\sin \varphi' = \frac{c + \sin \varphi}{1 + c \sin \varphi}, \quad \lambda' = \lambda$$



2. Nesting



3. Embedding of a high-resolution region



Test-driven development strategy

All model dynamics configuration are to pass the following test hierarchy (closely following DCMIP tests):

Test type	Typical problems	Tested model components
Shallow water tests	Solid rotation, Rossby-wave, Barotropic instability	Horizontal approximation
3D idealized tests (with and without orography)	Baroclinic instability, Mountain waves, etc.	+ vertical approximation, orography treatment, temporal splitting scheme (for NH)
Simple physics tests	Tropical cyclone, Held-Suarez	+ coupling subgrid-scale physics to dynamics
Real orography, real starting data test		+ accuracy and stability over complex terrain and complex flow conditions
Full physics simplified tests	Aqualanet	+ interaction of dynamics with complex subgrid-scale physics
Real forecast (climate) test		All



Test-driven development strategy

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Simple physics tests	Tropical cyclone, Held	coupling subgrid-scale physics to

The real (retrospective) forecast and historical climate simulations are final and ultimate model test. Right solutions are known (at least up to observations/(re)analysis accuracy).

But the problem is very complex and it is hard to find out what is the cause of deviations when something goes wrong

		complex subgrid-scale physics
Real forecast (climate) test		

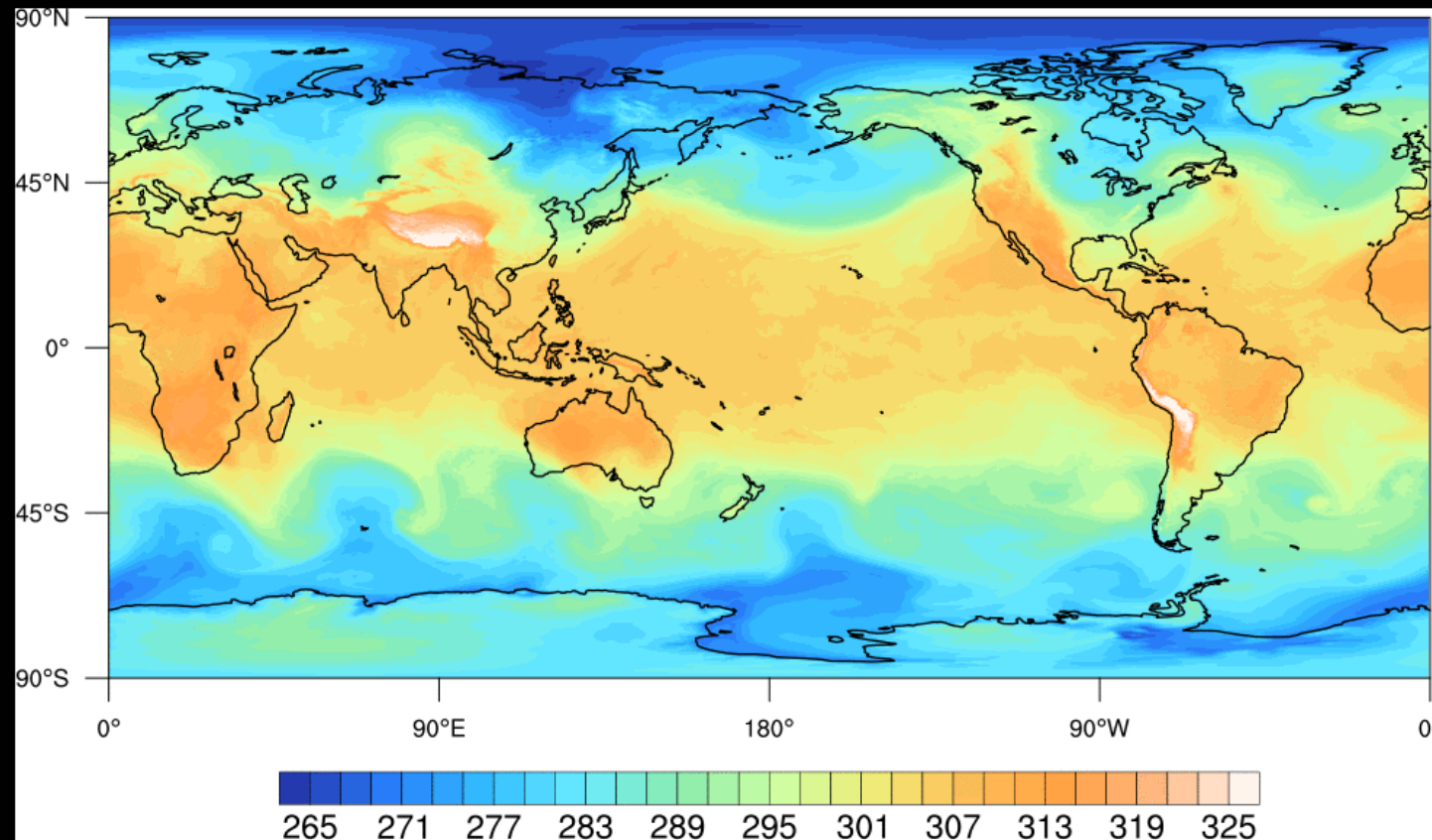


Real starting data



H

ons



0.25° r

d at level

Model stability in the absence of evident artifacts is primary results

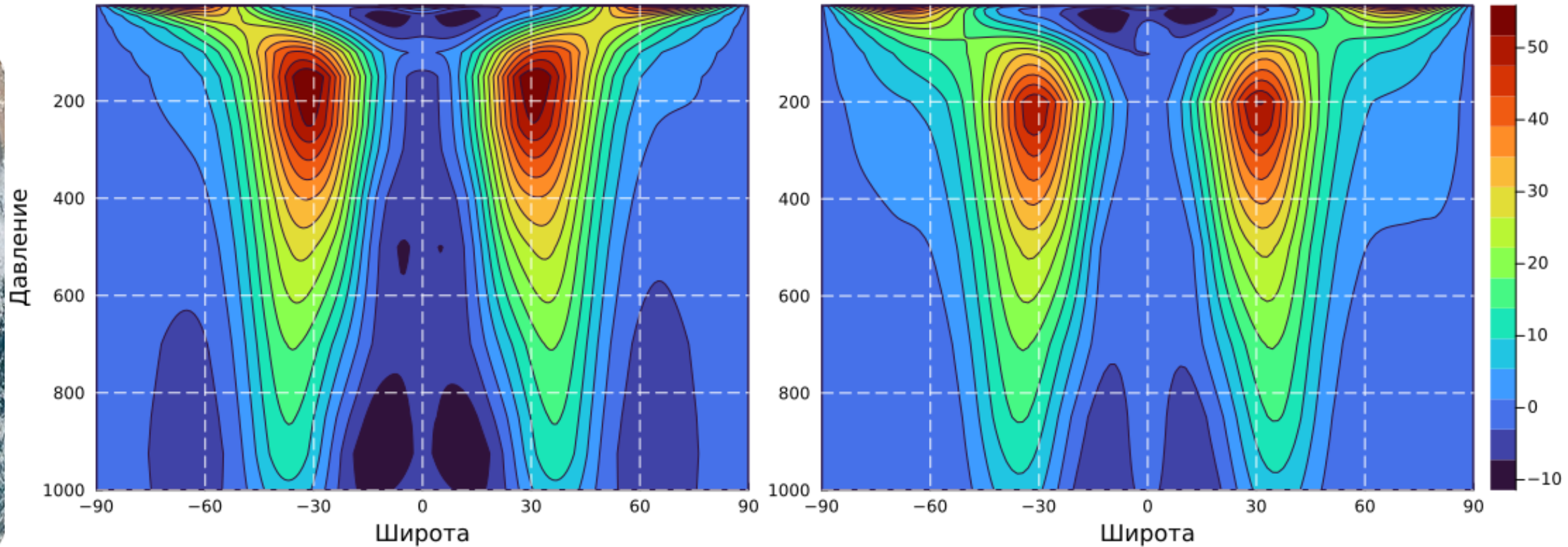
Aquaplanet experiment

- A planet covered by ocean, prescribed SST, ozone
- Diurnal cycle, perpetual equinox
- Full subgrid-scale physics package from SLAV model
- Isothermic initial conditions (6 months spin-up)
- Comparison with SLAV of comparable resolution of $0.96^\circ \times 0.72^\circ$ with the absolutely same subgrid scale physics package
- Cubed-sphere C90 (1°) grid, 96 vertical levels (SLAV vertical grid), dynamics $\text{dt}=180\text{s}$, physics $\text{dt} = 180\text{s}, 1800\text{s}$.
- 9 simulated years per day with 576 computational cores.

We are testing the very basic configuration of our dynamics in the very basic setup, but the first time with full subgrid-scale physics package

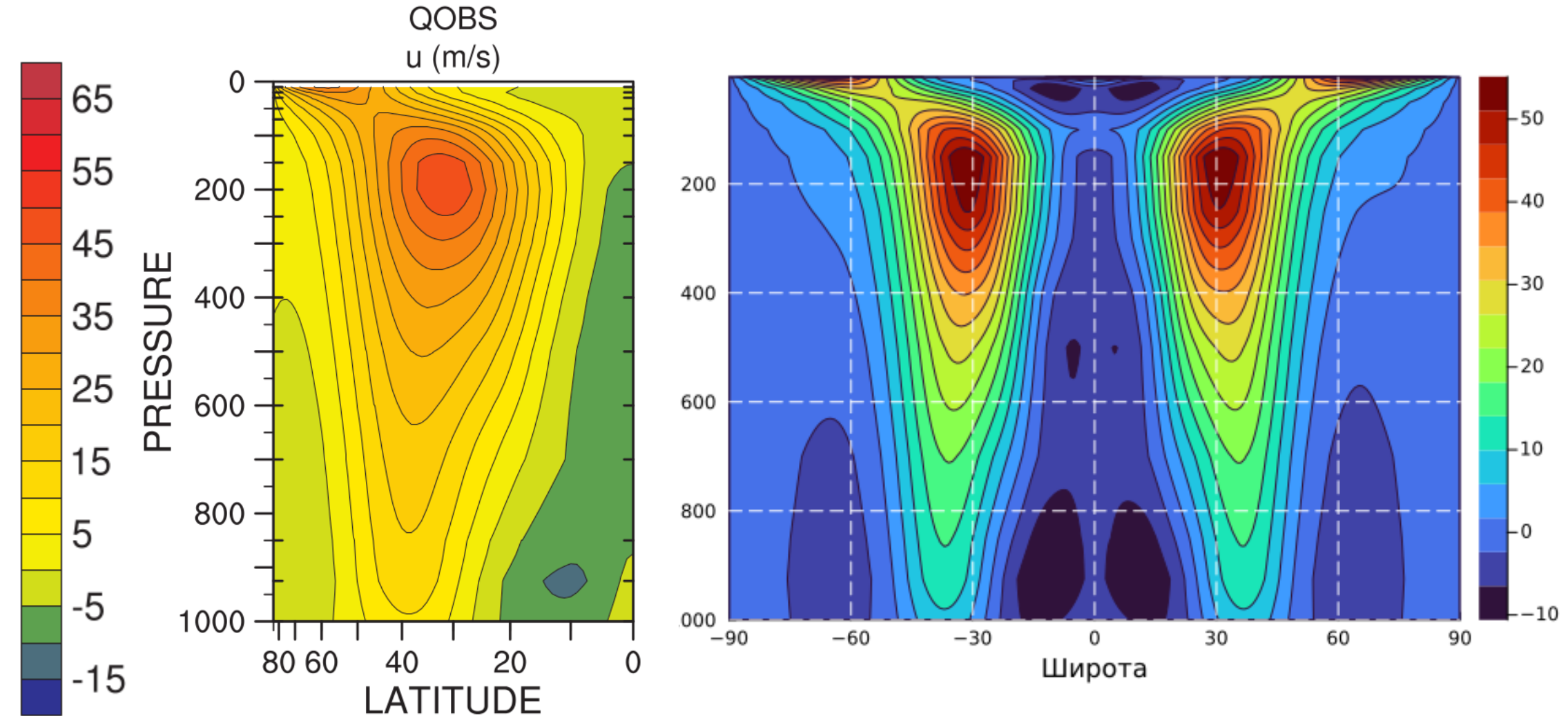


Aquaplanet experiment – zonal wind



2000 days average of zonal mean velocity according to cubed-sphere model (left) and SLAV (right)

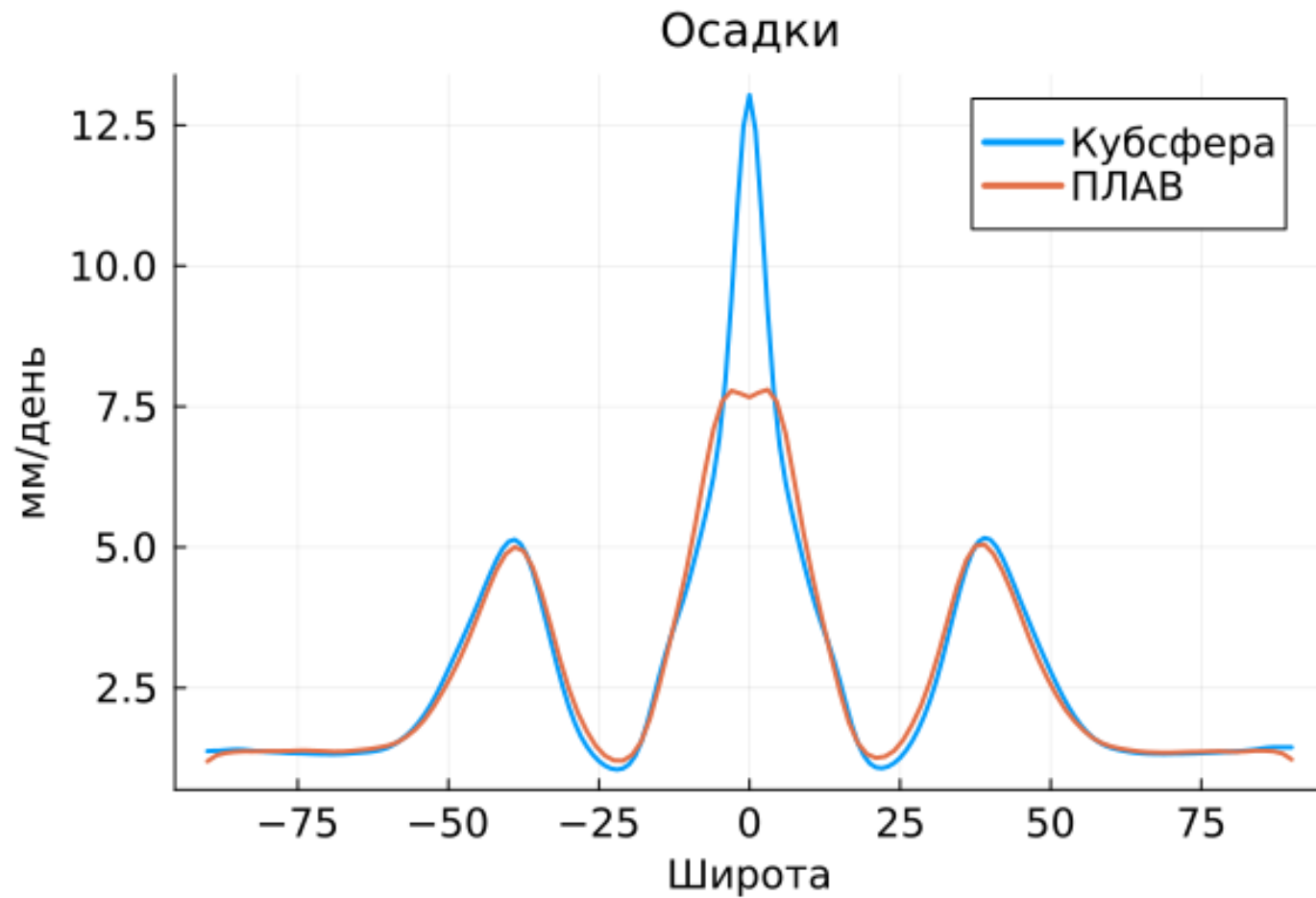
Aquaplanet experiment – zonal wind



2000 days average of zonal mean velocity. Left – multimodel mean, right – cubedsphere

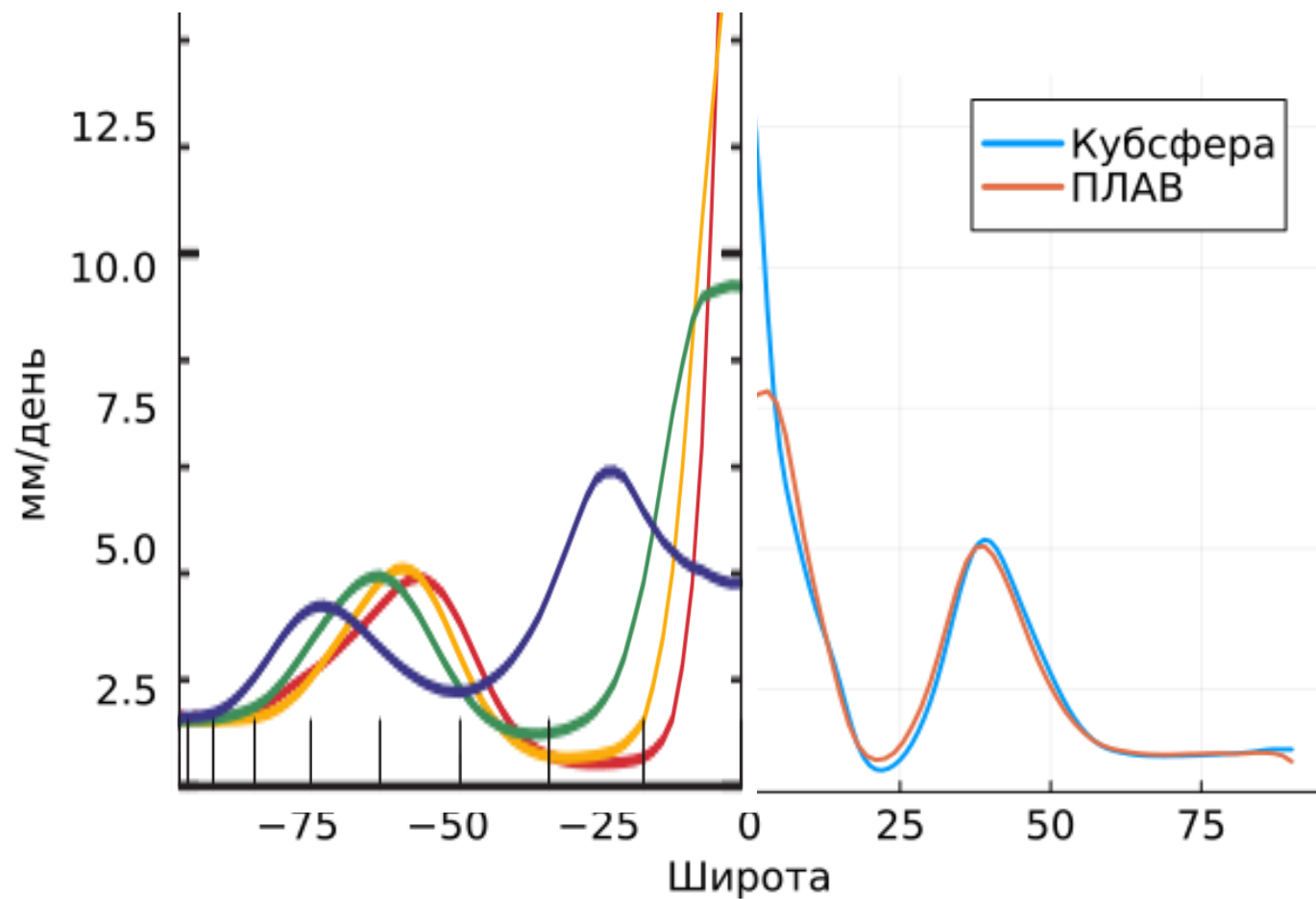


Aquaplanet experiment - precipitation



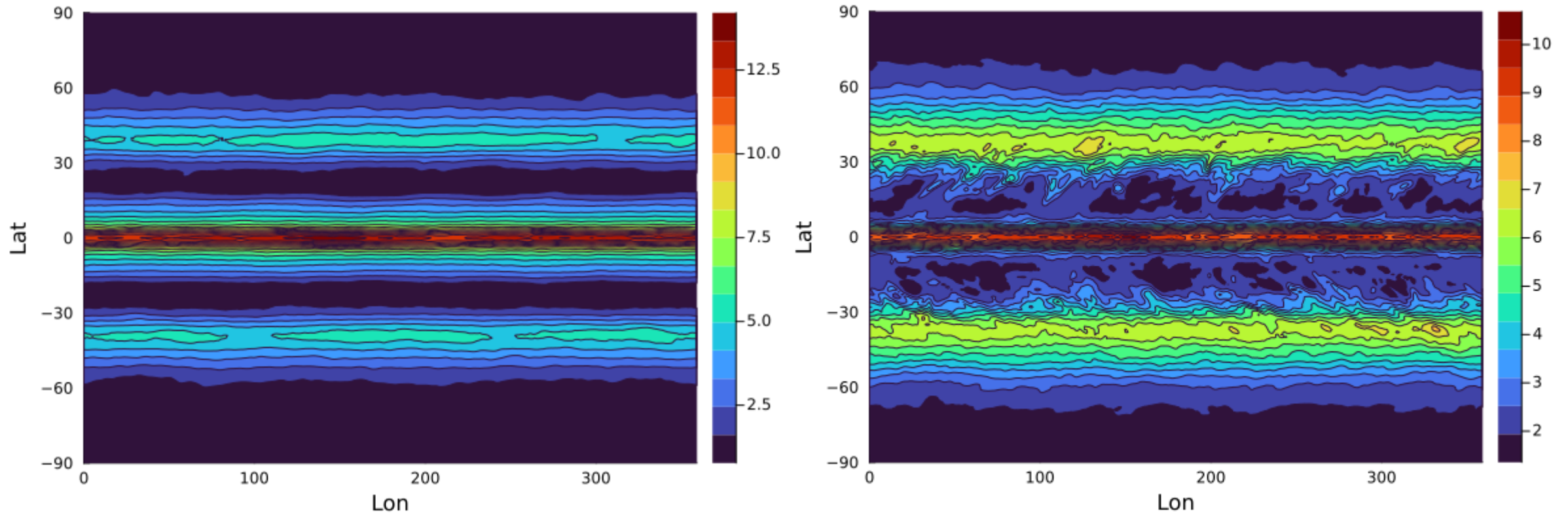
2000 days average of zonal mean total precipitation (blue – cubed-sphere, red - SLAV)

Aquaplanet experiment - precopotation



2000 days average of zonal mean total precipitation (blue – cubed-sphere, red - SLAV). Left – multimodel mean (green curve)

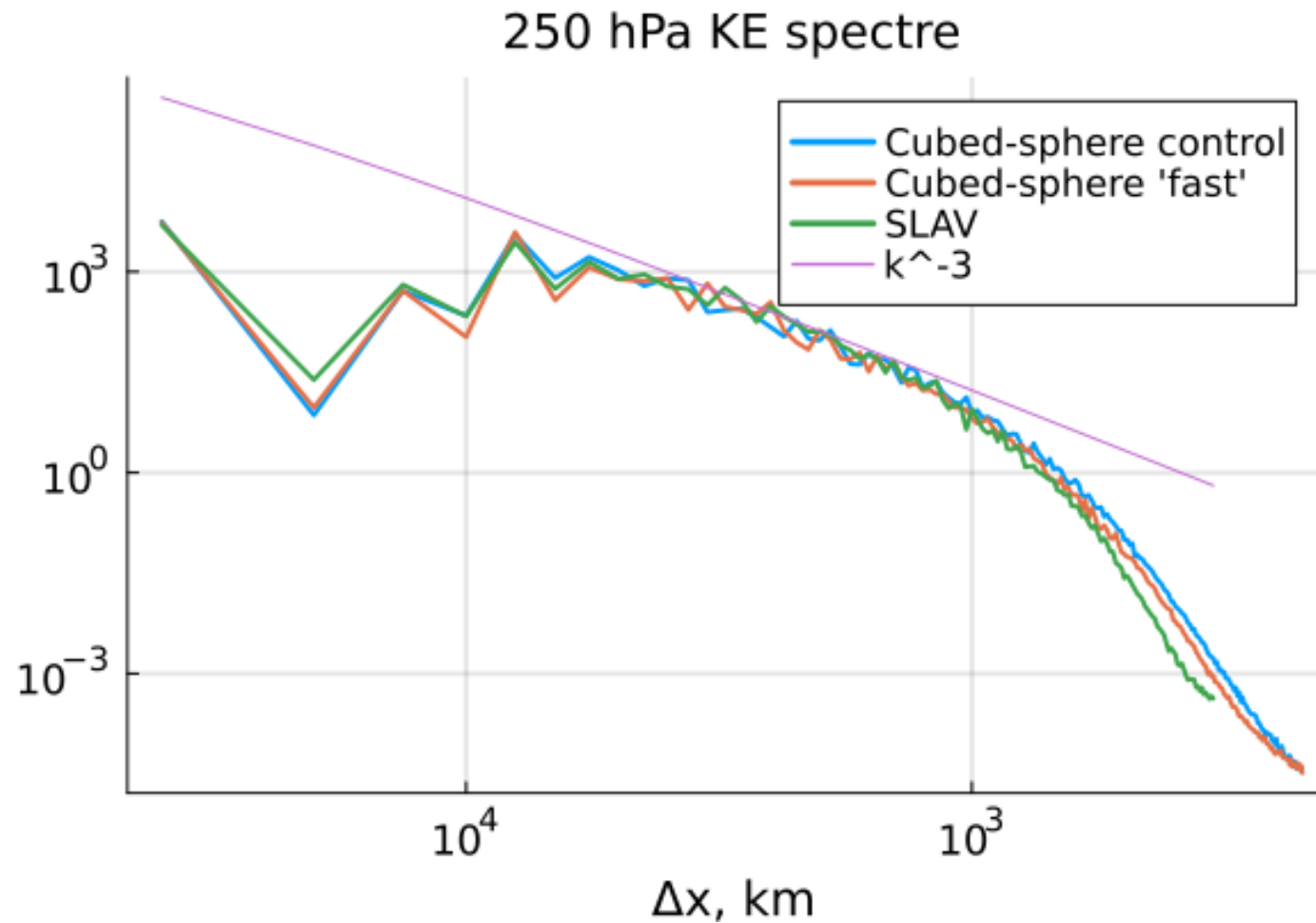
Aquaplanet experiment - precipitation



Left - 2000 days average total precipitation (mm/day). Right – 2000 days total precipitation standard deviation (mm/day). Both – cubed-sphere model.

No evident cubed-sphere grid imprinting (good)

Aquaplanet experiment - precipitation

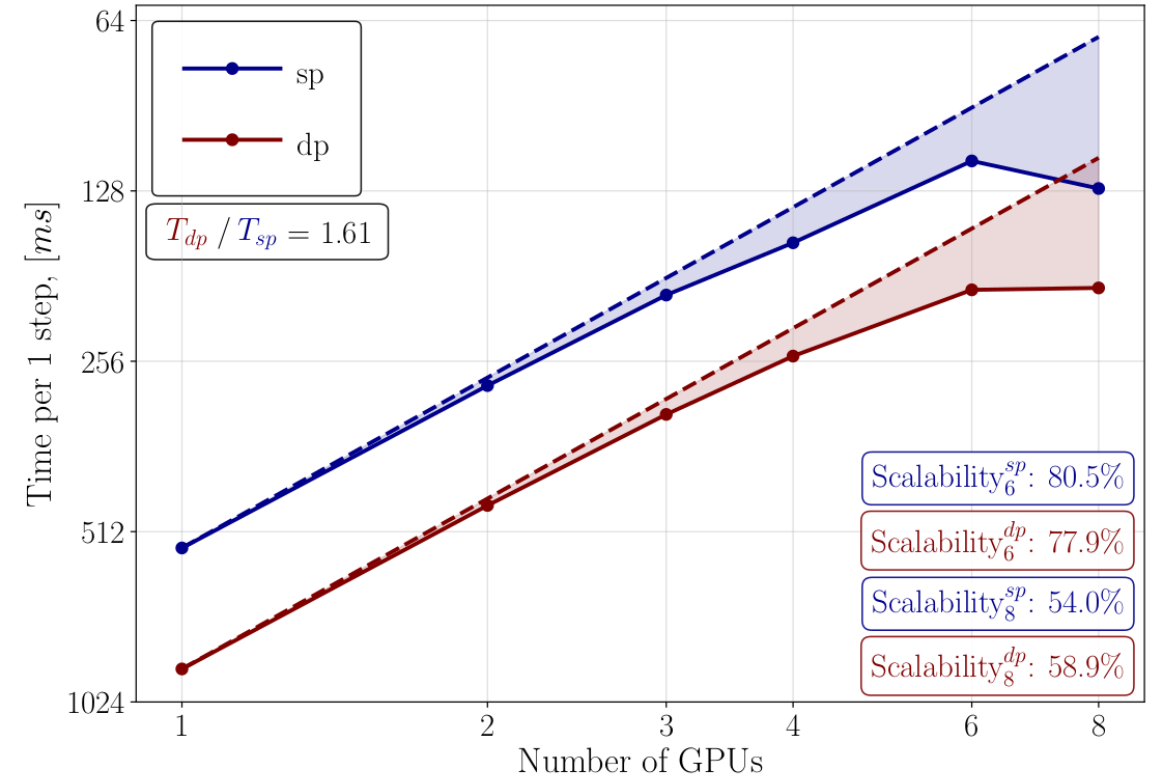
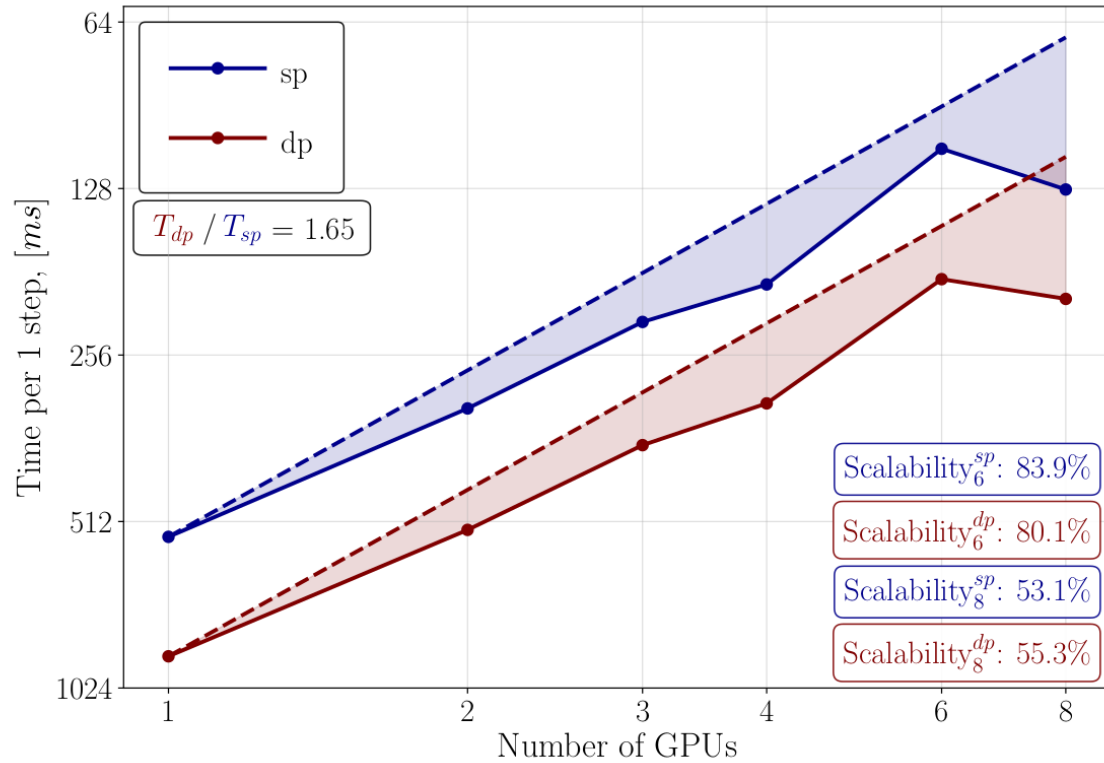


Kinetic energy spectre at 250 hPa level.

Cubed-sphere control: double precision in dynamics and physics each timestep

Cubed-sphere “fast”: single precision in dynamics, physics each 10-th dynamics timestep

A word on computational efficiency: GPU implementation performance



Strong scalability at RSK test GPU server (8 x H100). Left - standard, right using NVLink interconnect. Model configuration: C360 (0.25°), 96 levels. Dynamical core only (no subgrid-scale physics)

1 H100 $\approx$ 450 Xeon Gold CPU cores

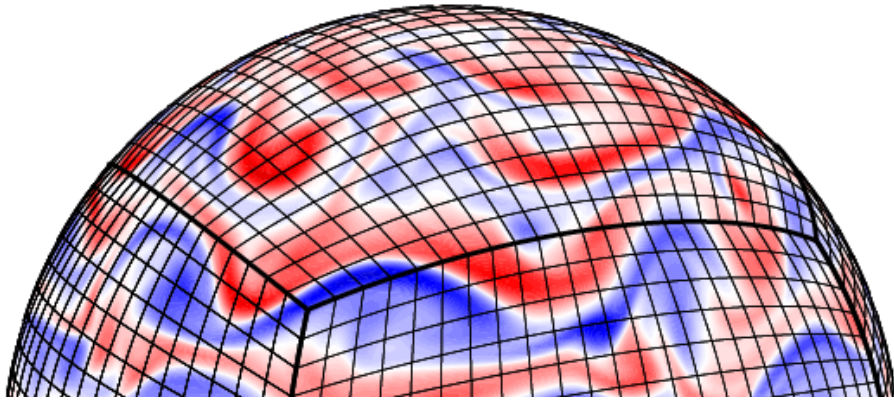
Conclusions

- On the way to real simulations with SLAV-model subgrid-scale physics (one step left)
- Yet to analyze and tune some things in aquaplanet experiment
- At the time results look promising

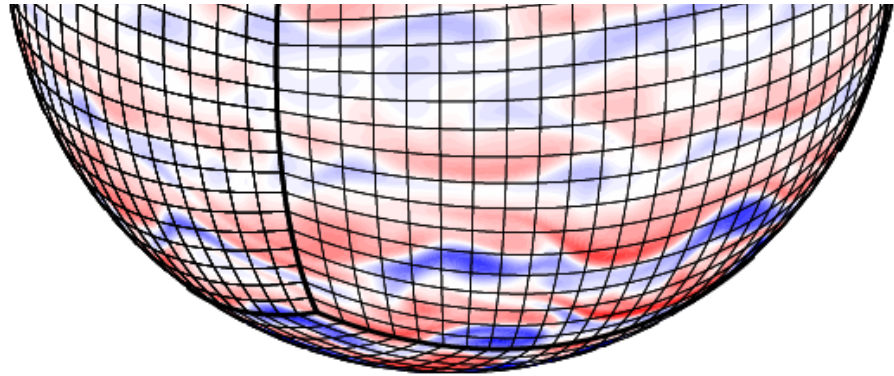
Outlook:

- Coupling with ARTS-ESM/INMCM physics
- Dynamical core code gives us much more opportunities that we can ever test (seeking for collaborations)
- A way to really high-resolution NWP and climate runs: making full-GPU code (GPU subgrid-scale physics)





Thank you for attention!



Vladimir Shashkin



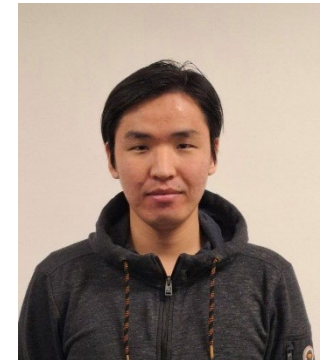
Gordey Goyman



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Dmitry Markhanov



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Special thanks to Mikhail Tolstykh and Rostislav Fadeev

Vladimir Shashkin (v.shashkin@inm.ras.ru)

<https://gitlab.inm.ras.ru/vshashkin/ParCS>



Non-hydrostatic equations

$$\frac{\partial \vec{v}_{hor}}{\partial t} = (\xi + f)k \times \vec{v}_{hor} - \nabla_{hor} K - C_p \theta \nabla_{hor} P, \quad (1a)$$

$$\frac{d \vec{v}_{hor}}{dt} = f k \times \vec{v}_{hor} - C_p \theta \nabla_{hor} P, \quad (1b)$$

$$\frac{dw}{dt} = -g - C_p \theta \frac{\partial P}{\partial z}, \quad (2)$$

$$\frac{d\theta}{dt} = 0, \quad (3a)$$

$$\frac{\partial \rho \theta}{\partial t} = - \left(\nabla_{hor} \cdot (\rho \theta \vec{v}_{hor}) + \frac{\partial \rho \theta w}{\partial z} \right), \quad (3b)$$

$$\frac{\partial \rho}{\partial t} = - \left(\nabla_{hor} \cdot (\rho \vec{v}_{hor}) + \frac{\partial \rho w}{\partial z} \right), \quad (4a)$$

$$\frac{d\rho}{dt} = - \nabla_{hor} \cdot \vec{v}_{hor} - \frac{\partial w}{\partial z}, \quad (4b)$$

$\vec{v}_{hor}, w$ - horizontal and vertical wind comp.

$$\xi = k \cdot \nabla \times \vec{v}_{hor}, \quad K = \left| \vec{v}_{hor} \right|^2$$

P - Exner pressure

θ - Potential temperature

ρ - air density

Basic solver– (1a), (2), (3a), (4a)

Energy-conserving – (1a), (2), (3b), (4a)

Semi-Lagrangian – (1b), (2), (3a), (4b)



Hydrostatic equations


$$\frac{\partial \vec{v}_{hor}}{\partial t} = (\zeta + f) \vec{v}_{hor}^{\perp} - \nabla_{hor} K - \nabla_{hor} \Phi - \frac{RT}{p} \frac{\partial p}{\partial p_s} \nabla_{hor} p_s,$$

$$\frac{\partial}{\partial t} \frac{\partial p}{\partial \eta} = -\nabla_{hor} \cdot \left(\frac{\partial p}{\partial \eta} \vec{v} \right) - \frac{\partial}{\partial \eta} \left(\frac{\partial p}{\partial \eta} \dot{\eta} \right),$$

$$\frac{\partial q}{\partial t} = -\vec{v} \cdot \nabla q - \dot{\eta} \frac{\partial q}{\partial \eta},$$

$$C_p \frac{dT}{dt} = \frac{RT}{p} \omega,$$

$$\frac{\partial \Phi}{\partial \eta} = -\frac{RT}{p} \frac{\partial p}{\partial \eta},$$

$$\zeta = k \cdot \nabla \times \vec{v}_{hor}, \quad K = \left| \vec{v}_{hor} \right|^2$$

$\vec{v}_{hor}$ - horizontal wind, $\vec{v}_{hor}^{\perp} = k \times \vec{v}_{hor}$

p - pressure, p_s - surf. pressure

η - vertical coordinate, $\dot{\eta} = \frac{d\eta}{dt}$, $\omega = \frac{dp}{dt}$

q - water vapour specific concentration,

Φ - geopotential, T - temperature

Скорость вычислений и масштабируемость на CPU

Cray XC40 (ГВЦ Росгидромета), Cray ARIES interconnect,
Intel Xeon E2697v4 18-core CPUs, 36 CPU на узле.

Конфигурации: **C90L73** (1^0 , 73 верт. уровня), **C360L73** (0.25^0 , 37 верт. уровня). Схема 2-го порядка, одинарная точность, эксперимент Хелда-Суареса.



C90L73

288 ядер – 41 лет/сутки
1152 ядра – 180 лет/сутки
1728 ядер – 237 лет/сутки

C360L37

1152 ядра – 5.2 лет/сутки, 26 дней/20 минут
6912 ядер – 26 лет/сутки, 130 дней/20 минут

Реальная подсеточная физика $\sim 2x$ замедление расчетов